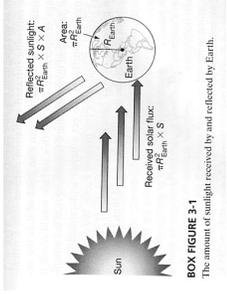




## Solar Energy Absorbed by Earth



**BOX FIGURE 3-1**  
The amount of sunlight received by and reflected by Earth.  
(from *The Earth System*)

- **Solar Constant (S)**  
= solar flux density reaching the Earth  
= 1370 W/m<sup>2</sup>
- **Solar energy incident on the Earth**  
= S x the “flat” area of the Earth  
= S x π R<sup>2</sup><sub>Earth</sub>
- **Solar energy absorbed by the Earth**  
= (received solar flux) – (reflected solar flux)  
= S π R<sup>2</sup><sub>Earth</sub> – S π R<sup>2</sup><sub>Earth</sub> x A  
= S π R<sup>2</sup><sub>Earth</sub> x (1-A)

A is the *planetary albedo* of the Earth, which is about 0.3.



ESS200A  
Prof. Jih-Yi Yu

## What Happens After the Earth Absorbs Solar Energy?

- The Earth warms up and has to emit radiative energy back to the space to reach a equilibrium condition.
- The radiation emitted by the Earth is called “terrestrial radiation” which is assumed to be like blackbody radiation.



ESS200A  
Prof. Jih-Yi Yu

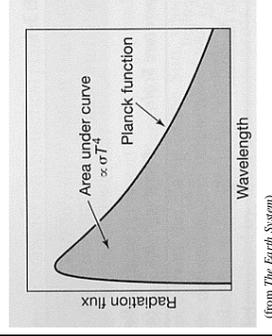
## Blackbody Radiation

- **Blackbody**  
A blackbody is something that emits (or absorbs) electromagnetic radiation with 100% efficiency at all wavelength.
- **Blackbody Radiation**  
The amount of the radiation emitted by a blackbody depends on the absolute temperature of the blackbody.



ESS200A  
Prof. Jih-Yi Yu

## Energy Emitted from Earth



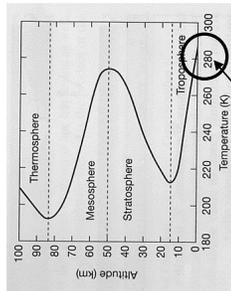
(from *The Earth System*)

- **The Stefan-Boltzmann Law**  
The energy flux emitted by a blackbody is related to the fourth power of the body's absolute temperature  
 $F = \sigma T^4$  where  $\sigma$  is  $5.67 \times 10^{-8} \text{ W/m}^2/\text{K}$
- **Energy emitted from the Earth**  
= (blackbody emission) x (total area of Earth)  
=  $(\sigma T_c^4) \times (4\pi R_{\text{Earth}}^2)$



ESS200A  
Prof. Jih-Yi Yu

# Planetary Energy Balance



(from Global Physical Climatology)

Energy emitted by Earth = Energy absorbed by Earth

$$\sigma T_e^4 \times (4\pi R_{\text{Earth}}^2) = S \pi R_{\text{Earth}}^2 (1-A)$$

$$\sigma T_e^4 = S/4 * (1-A)$$

$$= 1370/4 \text{ W/m}^2 * (1-A)$$

$$= 342.5 \text{ W/m}^2 * (1-A)$$

$$= 240 \text{ W/m}^2$$

Earth's blackbody temperature

$$T_e = 255 \text{ K } (-18\text{C})$$

Earth's surface temperature

$$T_s = 288 \text{ K } (15\text{C})$$

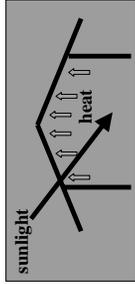
greenhouse effect (33C) !!



ESS200A  
Prof. Jih-Yi Yu

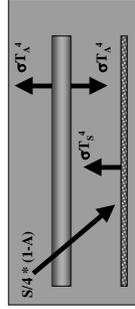
# Greenhouse Effect

## Greenhouse



- allow sunlight to come in
- trap heat inside the house

## Atmosphere



At the top of the atmosphere:

$$S/4 * (1-A) = \sigma T_a^4 \rightarrow T_a = T_e = 255\text{K}$$

For Earth's surface:

$$S/4 * (1-A) + \sigma T_a^4 = \sigma T_s^4$$

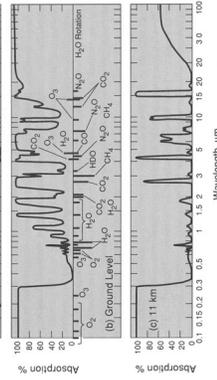
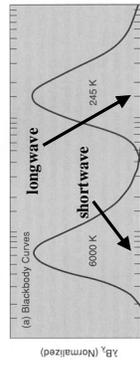
$$\rightarrow T_s = 1.19 T_a = 303\text{K}$$



ESS200A  
Prof. Jih-Yi Yu

# Different Wavelengths of Solar and Earth's Radiation

## Normalized Planck Function



(from Climate System Modeling)

## Planck Function

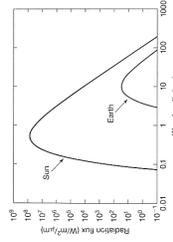


FIGURE 3-8 Planck curves for the Sun and Earth. The Sun emits most energy at short wavelengths.

(from The Earth System)



ESS200A  
Prof. Jih-Yi Yu

# Greenhouse Gases

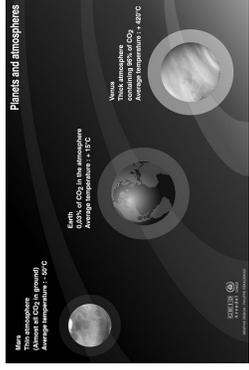
## Important Atmospheric Greenhouse Gases

| Name and Chemical Symbol                  | Concentration (ppm by volume)     |
|---|-----------------------------------|
| Water vapor, H <sub>2</sub> O             | 0.1 (South Pole)–40,000 (tropics) |
| Carbon dioxide, CO <sub>2</sub>           | 360                               |
| Methane, CH <sub>4</sub>                  | 1.7                               |
| Nitrous oxide, N <sub>2</sub> O           | 0.3                               |
| Ozone, O <sub>3</sub>                     | 0.01 (at the surface)             |
| Freon-11, CCl <sub>3</sub> F              | 0.00026                           |
| Freon-12, CCl <sub>2</sub> F <sub>2</sub> | 0.00047                           |



ESS200A  
Prof. Jih-Yi Yu

# Earth, Mars, and Venus



| Planet | Distance to the Sun | Radius    | Planetary Albedo | Mean Surface Temperature |
|--------|---------------------|-----------|------------------|--------------------------|
| Venus  | 0.72 AU             | 12,104 km | 0.80             | 730°K                    |
| Earth  | 1.00 AU             | 6,370 km  | 0.30             | 288°K                    |
| Mars   | 1.52 AU             | 6,794 km  | 0.22             | 218°K                    |

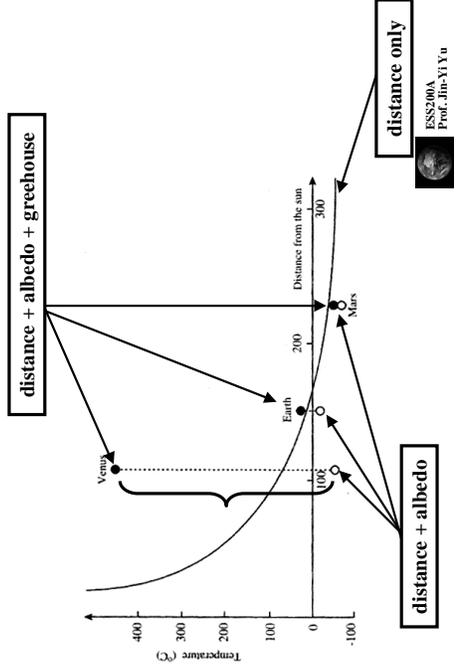
ESS200A  
Prof. Jih-Yi Yu

# Three Factors To Determine Planet Temperature

- Distance from the Sun
- Albedo
- Greenhouse effect

ESS200A  
Prof. Jih-Yi Yu

# Global Temperature



ESS200A  
Prof. Jih-Yi Yu

# Greenhouse Effects

- On Venus → 510°K (very large!!)
- On Earth → 33°K
- On Mars → 6°K (very small)

ESS200A  
Prof. Jih-Yi Yu

## Why Large Greenhouse Effect On Venus?

- **Venus is too close to the Sun**
- Venus temperature is very high
- Very difficult for Venus' s atmosphere to get saturated in water vapor
- Evaporation keep on bringing water vapor into Venus' s atmosphere
- Greenhouse effect is very large
- A "run away" greenhouse happened on Venus
- Water vapor is dissociated into hydrogen and oxygen
- Hydrogen then escaped to space and oxygen reacted with carbon to form carbon dioxide
- **No water left on Venus (and no more chemical weathering)**



ESS200A  
Prof. Jih-Yi Yu

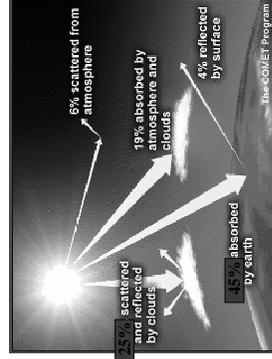
## Why Small Greenhouse Effect on Mars?

- **Mars is too small in size**
- Mars had no large internal heat
- Mars lost all the internal heat quickly
- No tectonic activity on Mars
- Carbon can not be injected back to the atmosphere
- Little greenhouse effect
- **A very cold Mars!!**



ESS200A  
Prof. Jih-Yi Yu

## Where Does the Solar Energy Go?



Incoming solar energy (100)

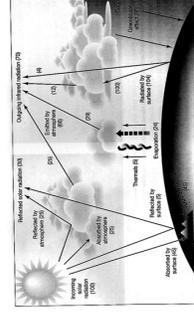
- **70% absorbed**
  - 45% by Earth' s surface (ocean + land)
  - 25% by the atmosphere and clouds
- **30% reflected and scattered back**
  - 20% by clouds
  - 6% by the atmosphere
  - 4% by surface



ESS200A  
Prof. Jih-Yi Yu

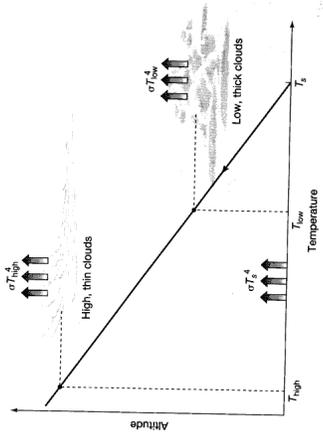
## Where Is Earth' s Radiation Emitted From?

- Radiation back to Space (70 Units)
- **70 (units) radiation back to space**
    - 66% by the atmosphere
    - 4% by surface (through clear sky)
  - **Greenhouse emission (back to surface)**
    - 88% (of solar radiation)



ESS200A  
Prof. Jih-Yi Yu

# Important Roles of Clouds In Global Climate



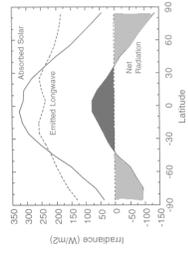
(from The Earth System)

ESS200A  
Prof. Jin-Yi Yu

ESS200A  
Prof. Jin-Yi Yu

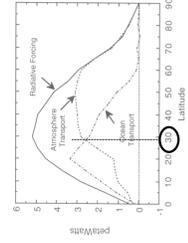
# Polarward Energy Transport

Annual-Mean Radiative Energy



Polarward heat flux is needed to transport radiative energy from the tropics to higher latitudes

Polarward Heat Flux



The atmosphere dominates the polarward heat transport at middle and high latitudes. The ocean dominates the transport at lower latitudes.

(1 peta Watts =  $10^{15}$  W)

(figures from Global Physical Climatology)

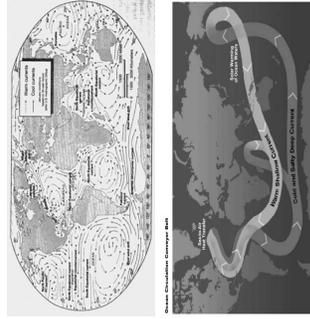
# How Do Atmosphere and Ocean Transport Heat?

Atmospheric Circulation



(from USA Today)

Ocean Circulation



(top from The Earth System)

(bottom from USGCRP)

ESS200A  
Prof. Jin-Yi Yu