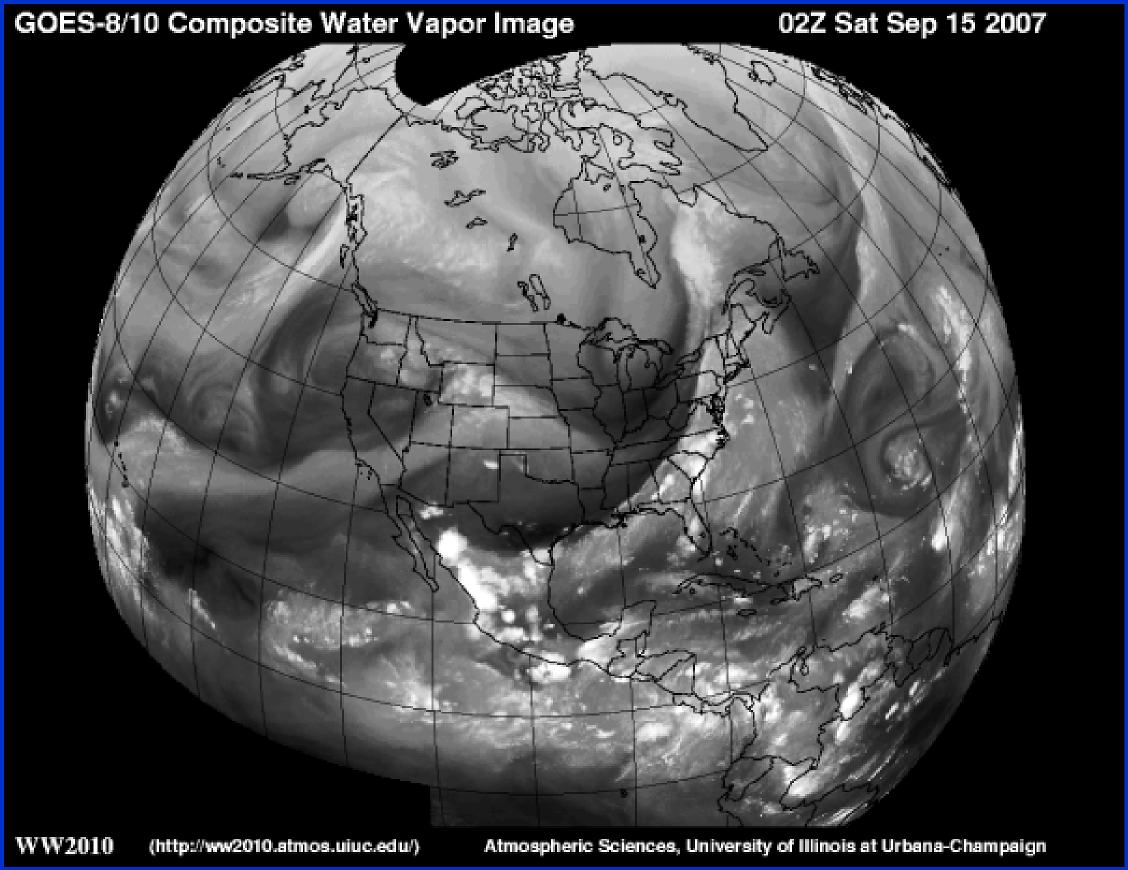
Chapter 7: Forces and Force Balances



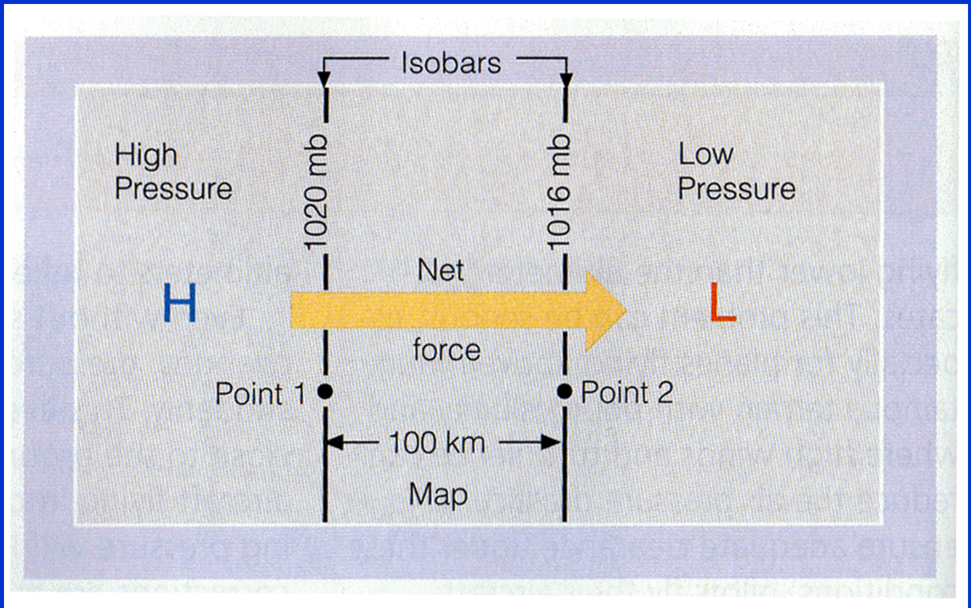
- Forces that Affect Atmospheric Motion
- Force Balance
- Geostrophic Balance and Jetstream

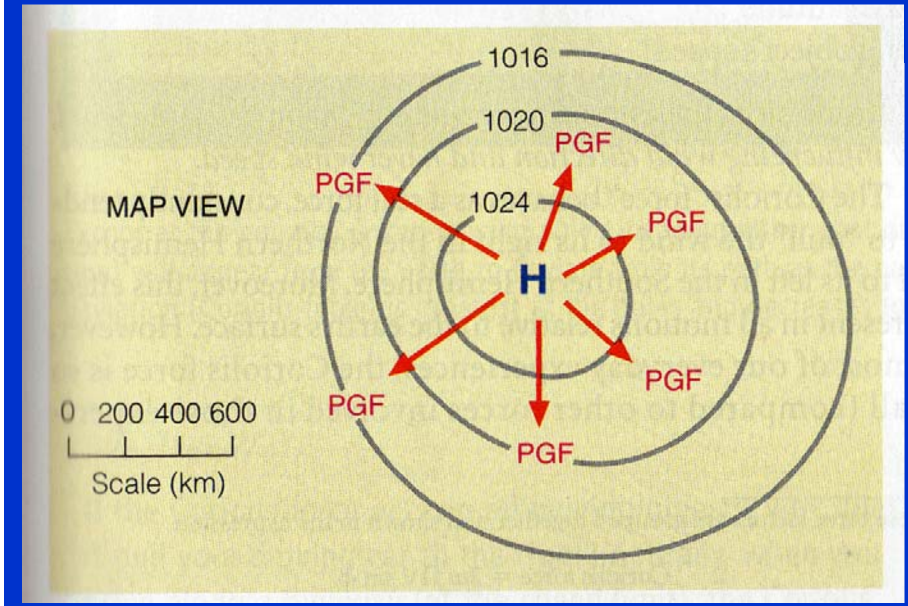
Forces that Affect Atmospheric Motion

- Fundamental force -
  - Pressure gradient force
  - Gravitational force
  - Frictional force
- Apparent force -
  - Centrifugal force
  - Coriolis force

- Newton's second law of motion states that the rate of change of momentum (i.e., the acceleration) of an object, as measured relative to coordinates fixed in space, equals the sum of all the forces acting.
- For atmospheric motions of meteorological interest, the forces that are of primary concern are the pressure gradient force, the gravitational force, and friction. These are the *fundamental* forces.
- For a coordinate system rotating with the earth, Newton's second law may still be applied provided that certain *apparent* forces, the centrifugal force and the Coriolis force, are included among the forces acting.

Pressure Gradient Force



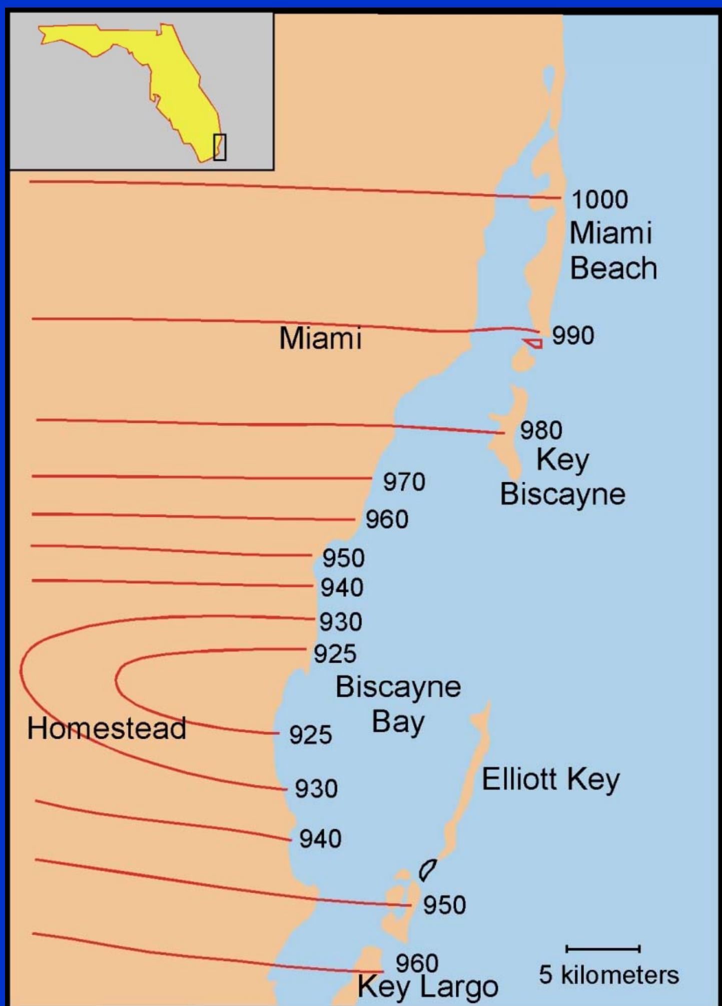


(from Meteorology Today)

- PG = (pressure difference) / distance
- Pressure gradient force goes from high pressure to low pressure.
- Closely spaced isobars on a weather map indicate steep pressure gradient.

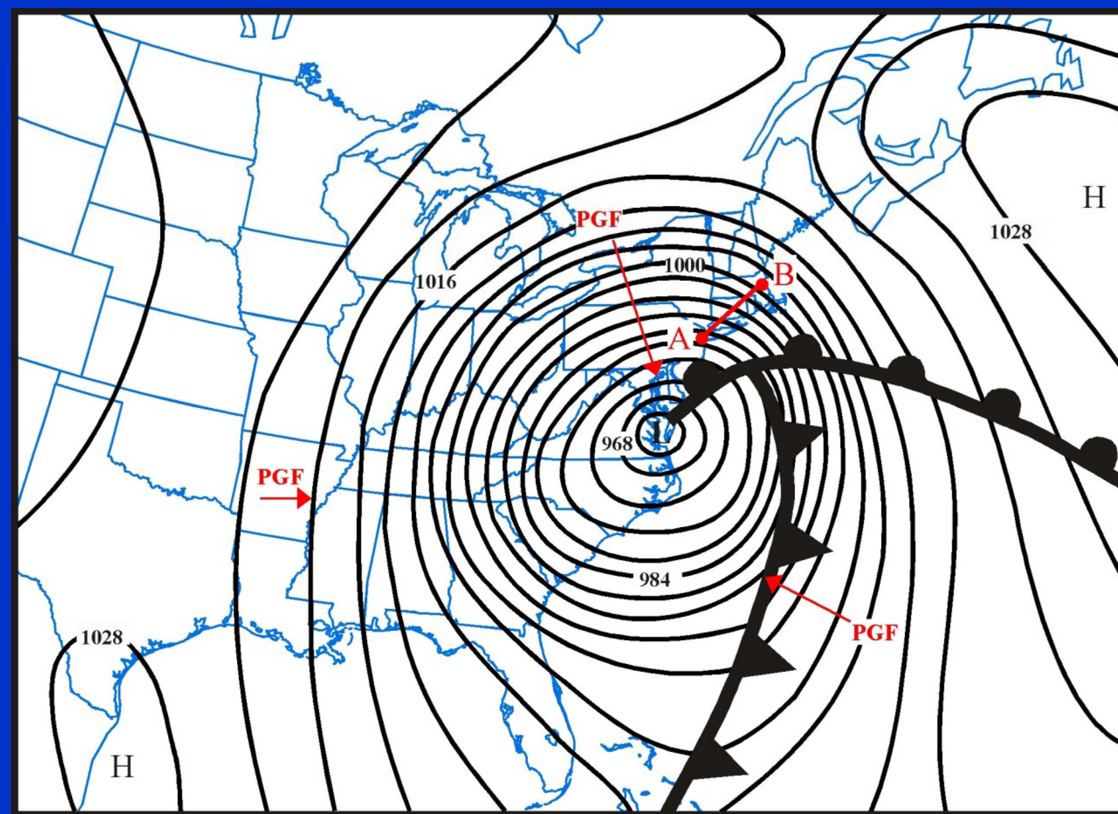
Examples of Pressure Gradient

Hurricane Andrew, 1992



Courtesy of American Meteorological Society

Extratropical Cyclone

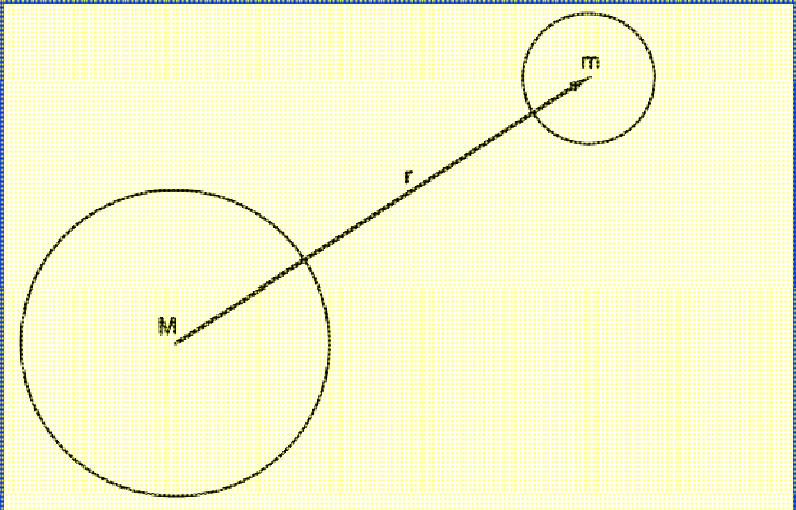


Pressure Gradients

- Pressure Gradients
  - The pressure gradient force initiates movement of atmospheric mass, wind, from areas of higher to areas of lower pressure
- Horizontal Pressure Gradients
  - Typically only small gradients exist across large spatial scales (1mb/100km)
  - Smaller scale weather features, such as hurricanes and tornadoes, display larger pressure gradients across small areas (1mb/6km)
- Vertical Pressure Gradients
  - ***Average vertical pressure gradients are usually greater than extreme examples of horizontal pressure gradients*** as pressure always decreases with altitude (1mb/10m)

Gravitational Force

- Newton's law of universal gravitation states that any two elements of mass in the universe attract each other with a force proportional to their masses and inversely proportional to the square of the distance separating them.
- Thus, if the earth is designated as mass M and m is a mass element of the atmosphere, then the force per unit mass exerted on the atmosphere by the gravitational attraction of the earth is



Two spherical masses whose centers are separated by a distance r.

$$\frac{\mathbf{F}_g}{m} \equiv \mathbf{g}^* = -\frac{GM}{r^2}\left(\frac{\mathbf{r}}{r}\right)$$

$r = a + z$ (a: earth radius; z: height above surface)

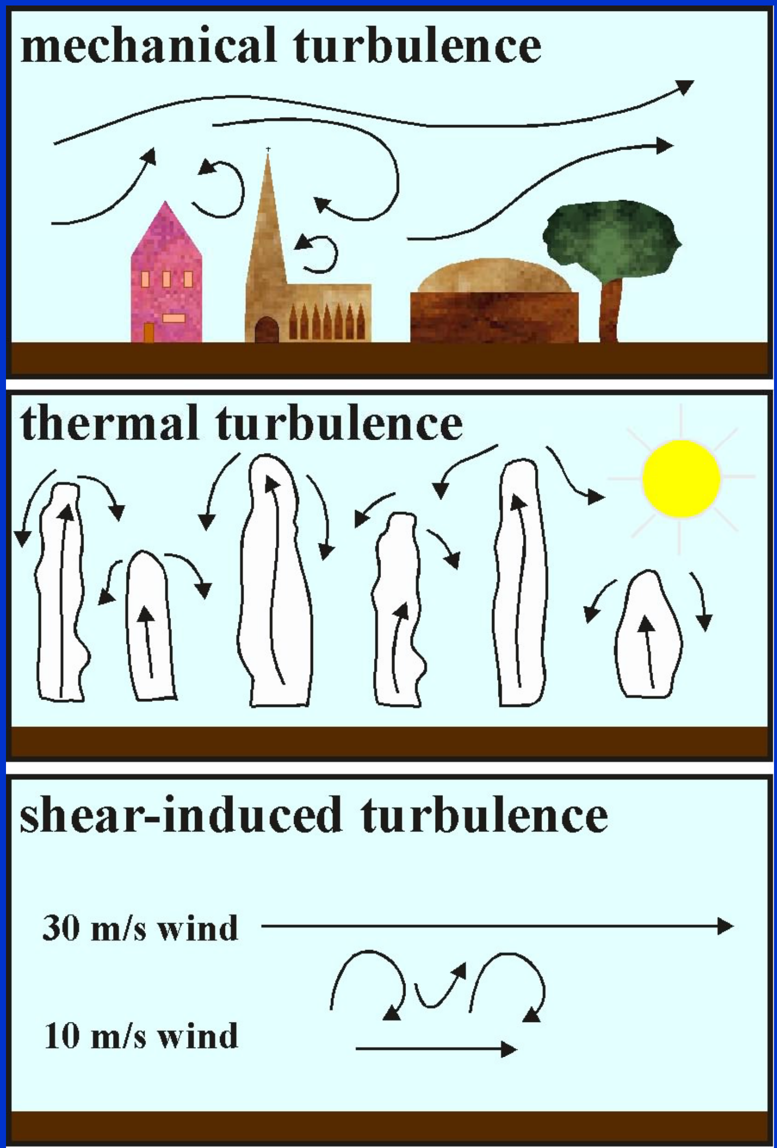
$$\mathbf{g}^* = \frac{\mathbf{g}_0^*}{(1 + z/a)^2}$$

where $\mathbf{g}_0^* = -(GM/a^2)(\mathbf{r}/r)$ is the gravitational force at mean sea level.

For meteorological applications,

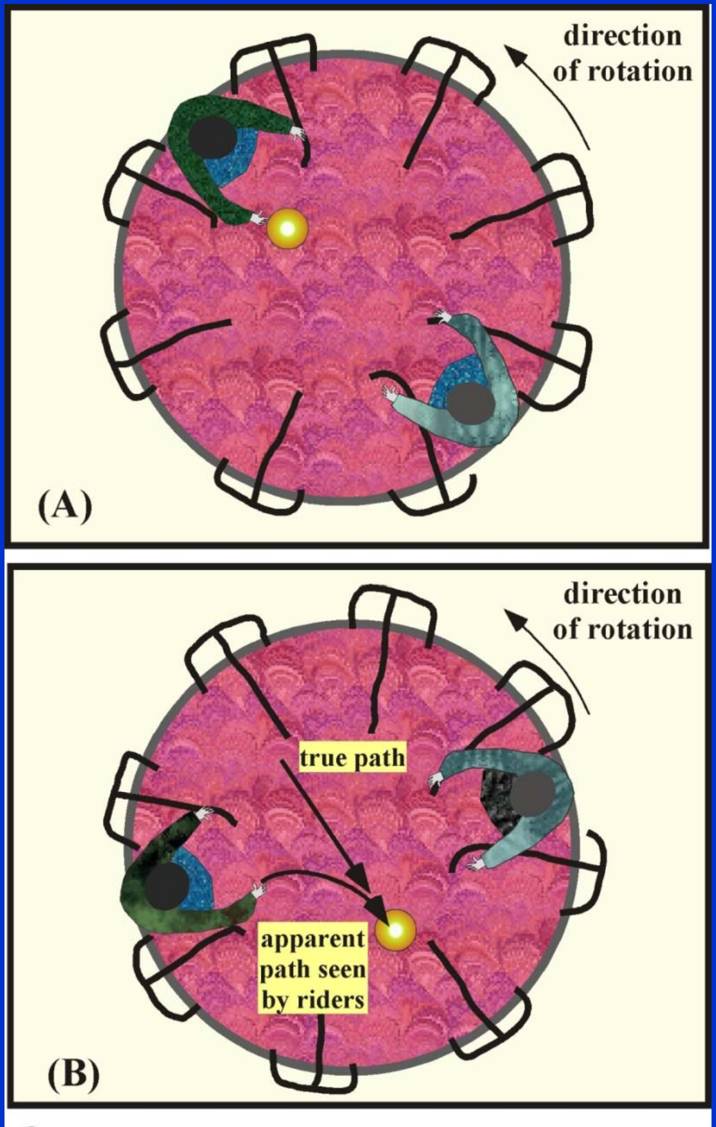
$z \ll a$ ➔ $\mathbf{g}^* = \mathbf{g}_0^*$

Frictional Force

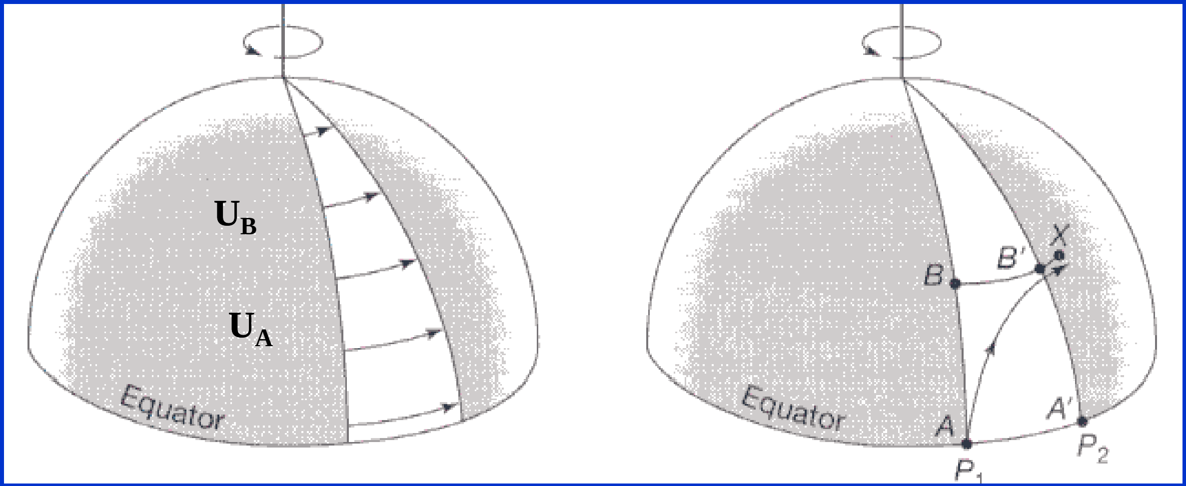


- Frictional force (drag) is strongest near the Earth's surface and decreases rapidly with height.
- The atmospheric layer in which frictional force is important is call thed boundary layer, whose depth can vary from a few hundred meters to a few thousand meters.
- There are three sources to generate turbulence eddies to give rise to the frictional force: (1) mechanical turbulence (airs encounter surface roughness), (2) thermal turbulence (air near Earth's surface get heated, and (3) wind-shear induced turbulence.

Example on a Merry-Go-Around



Coriolis Force



(from *The Earth System*)

- First, Point A rotates faster than Point B ($U_A > U_B$)

➔ $U_A > U_B$

➔ A northward motion starting at A will arrive to the east of B

➔ It looks like there is a "force" pushing the northward motion toward right

➔ This apparent force is called "Coriolis force":

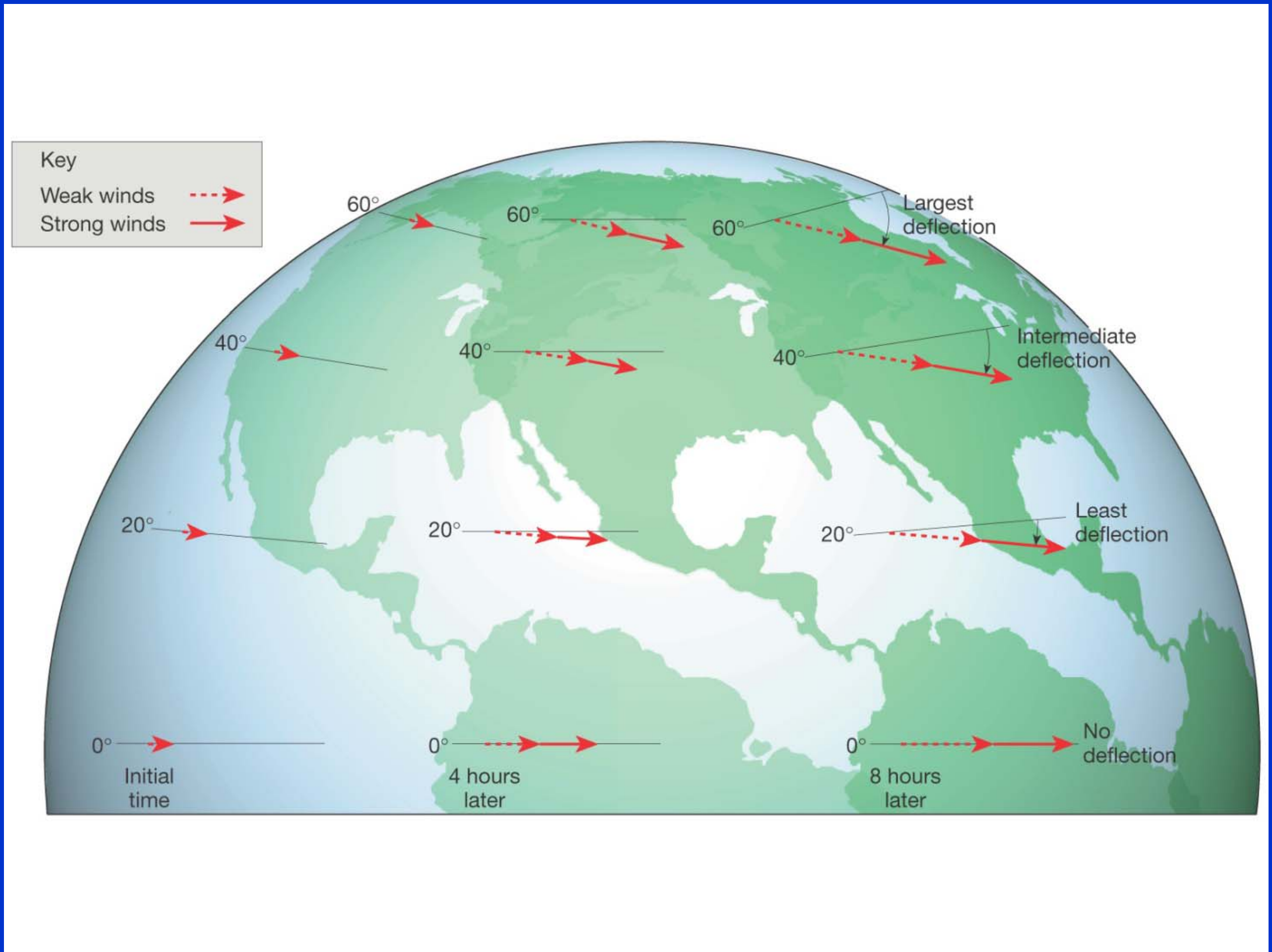
Coriolis Force = f V

where $f = 2*\Omega*\mathrm{Sin(lat)}$ and $\Omega = 7.292\times10^{-5}$ rad s^{-1}

Coriolis Force

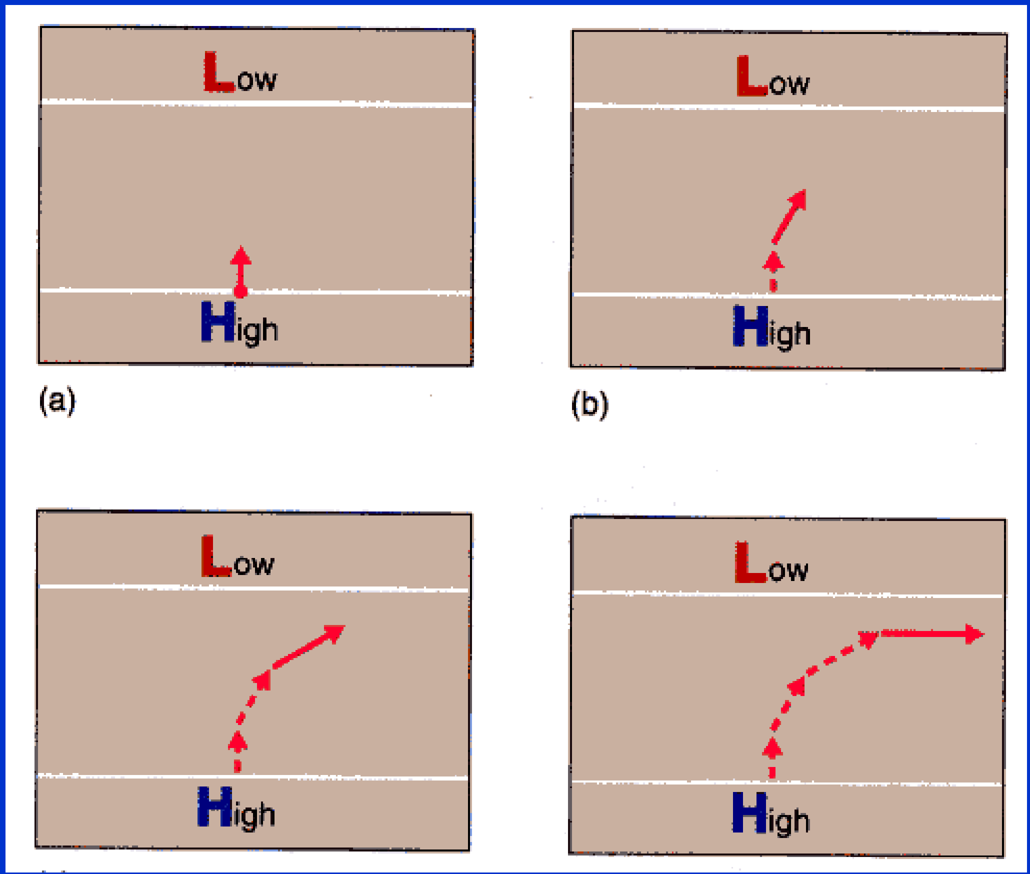
- Coriolis force causes the wind to deflect to the right of its intent path in the Northern Hemisphere and to the left in the Southern Hemisphere.
- The magnitude of Coriolis force depends on (1) the rotation of the Earth, (2) the speed of the moving object, and (3) its latitudinal location.
- The stronger the speed (such as wind speed), the stronger the Coriolis force.
- The higher the latitude, the stronger the Coriolis force.
- The Corioils force is zero at the equator.
- Coriolis force is one major factor that determine weather pattern.

Coriolis Force Change with latitudes



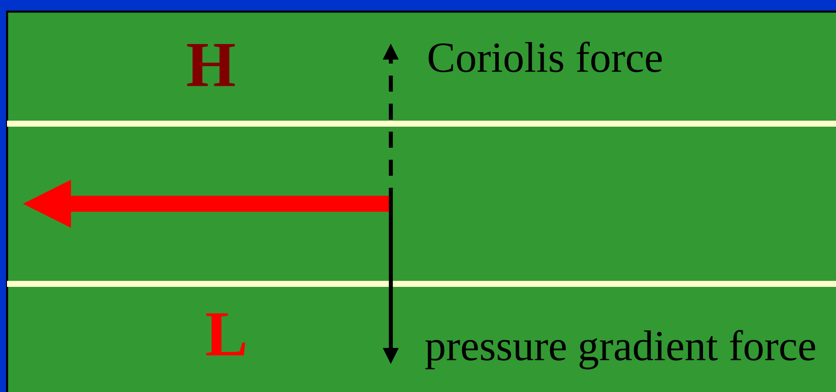
(from *The Atmosphere*)

How Does Coriolis Force Affect Wind Motion?



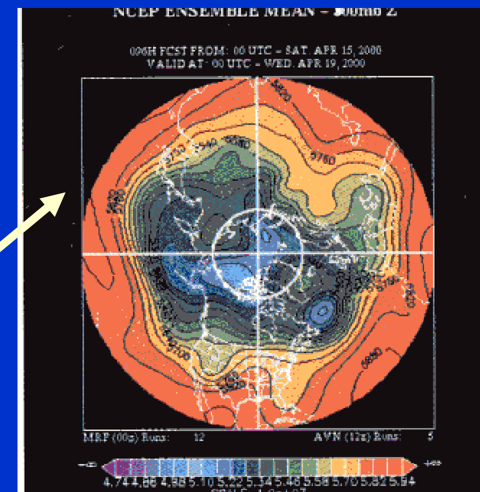
(from *Weather & Climate*)

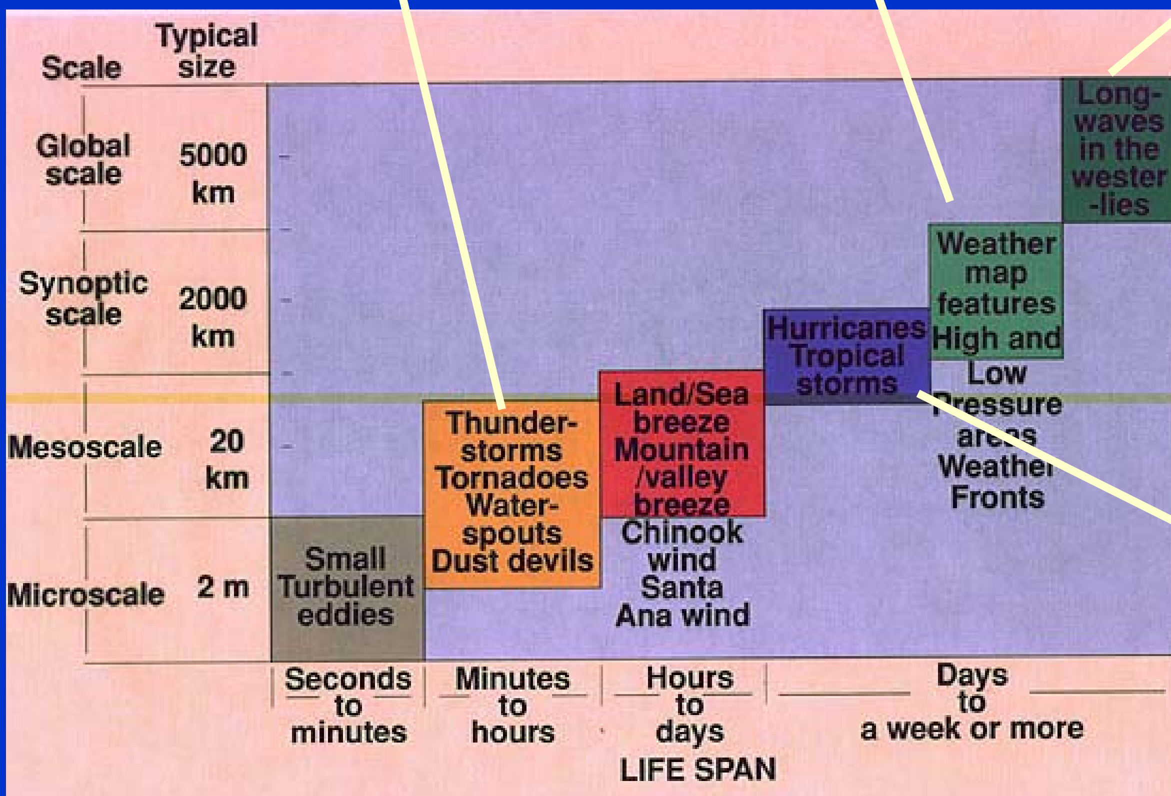
Geostrophic Balance



- By doing scale analysis, it has been shown that large-scale and synoptic-scale weather system are in geostropic balance.
- Geostrophic winds always follow the constant pressure lines (isobar). Therefore, we can figure out flow motion by looking at the pressure distribution.

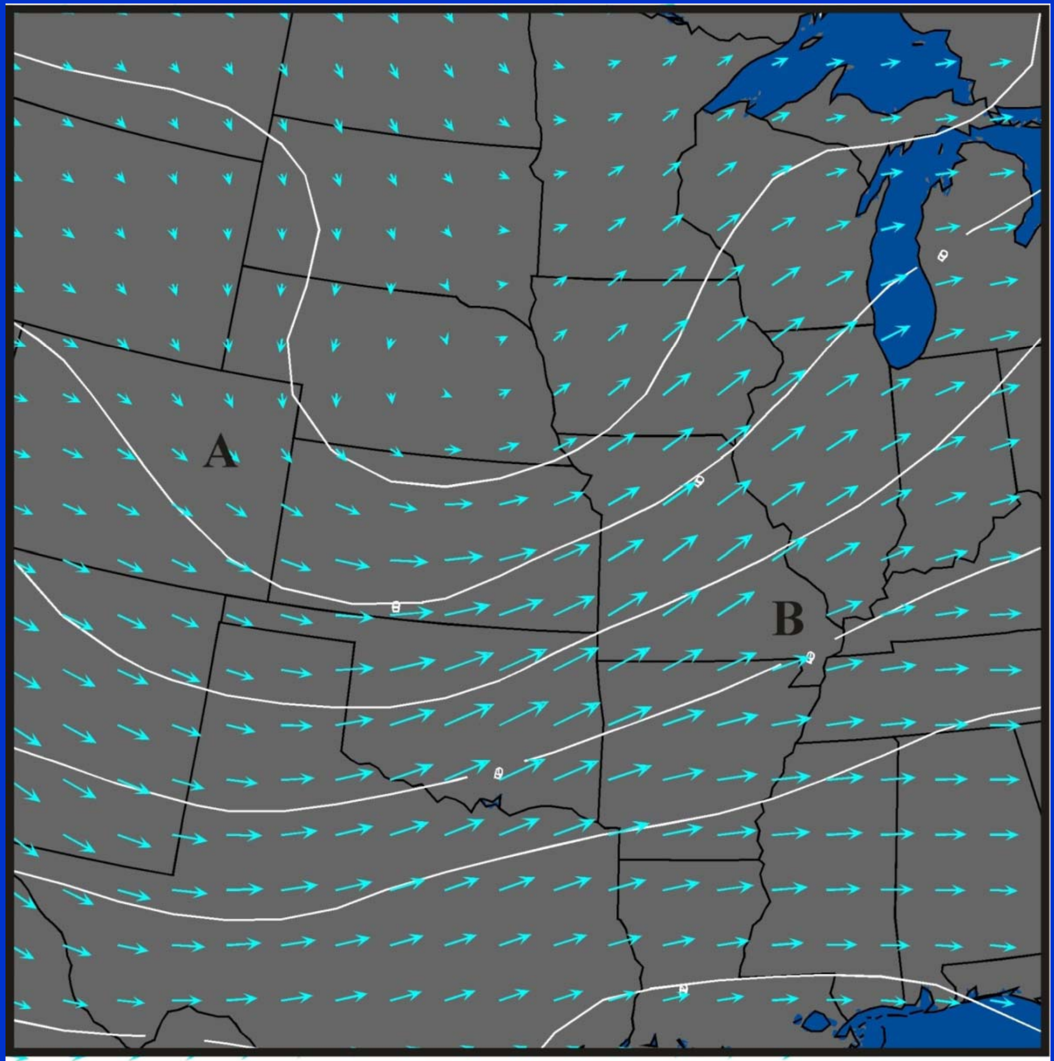
Scales of Motions in the Atmosphere





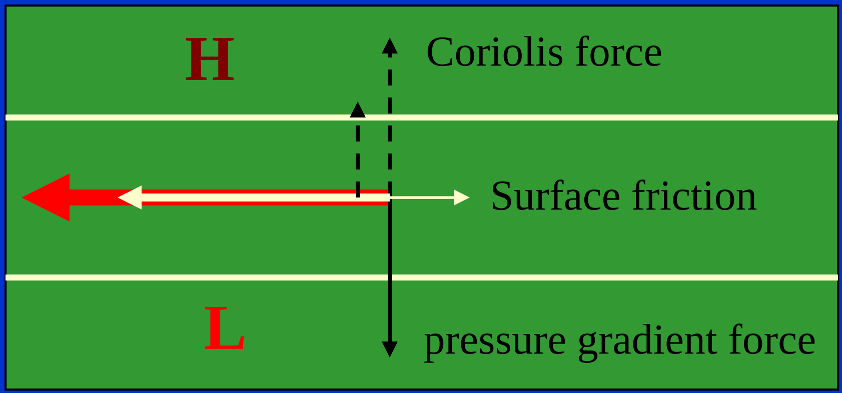
(from Meteorology Today by C. Donald Ahrens © 1994 West Publishing Company)

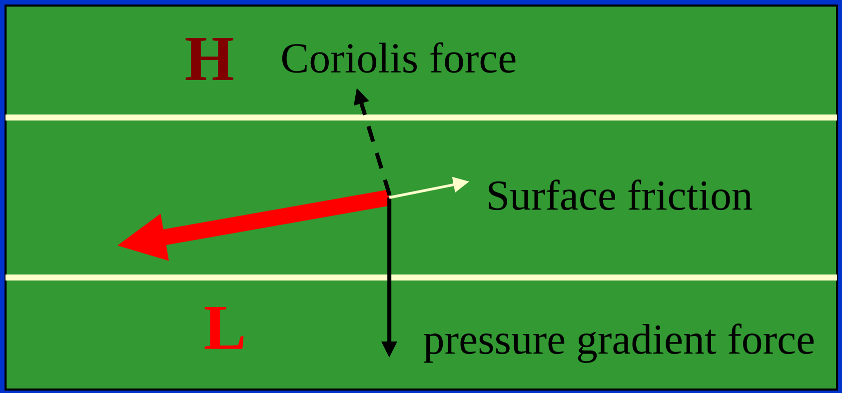
Example: Winds and Height on 500mb



Courtesy of the Department of Atmospheric Sciences
University of Illinois at Urbana-Champaign

Frictional Effect on Surface Flow





- Surface friction force slows down the geostrophic flow.
- The flow turns into (out of) the low (high) press sides.
- Convergence (divergence) is produced with the flow.

Surface Friction

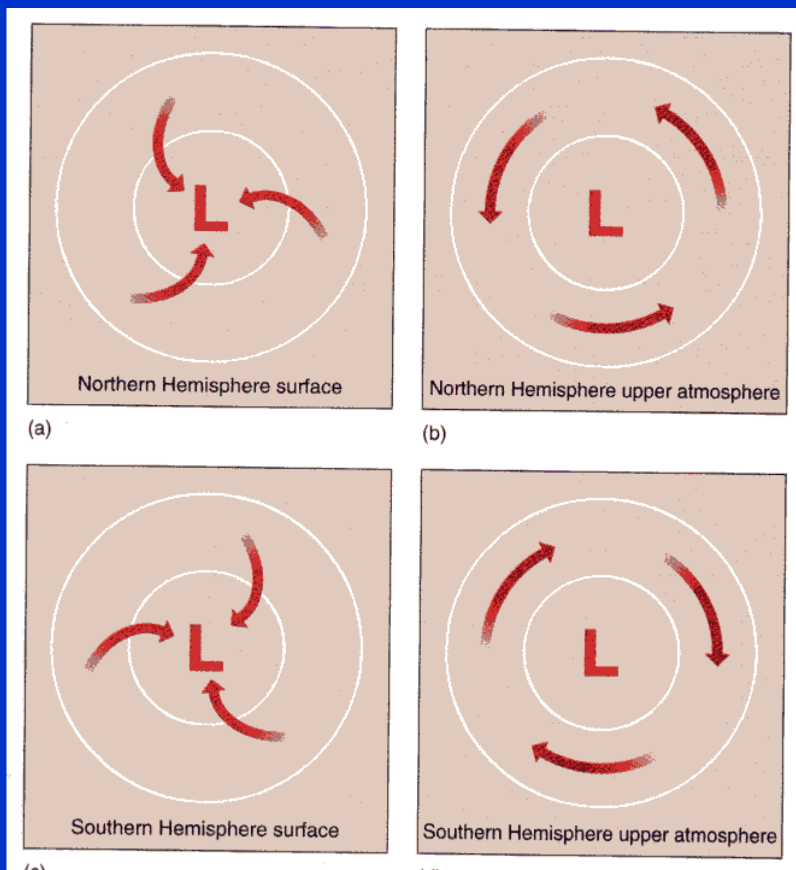
- Friction Force = $c * V$

  c = friction coefficient

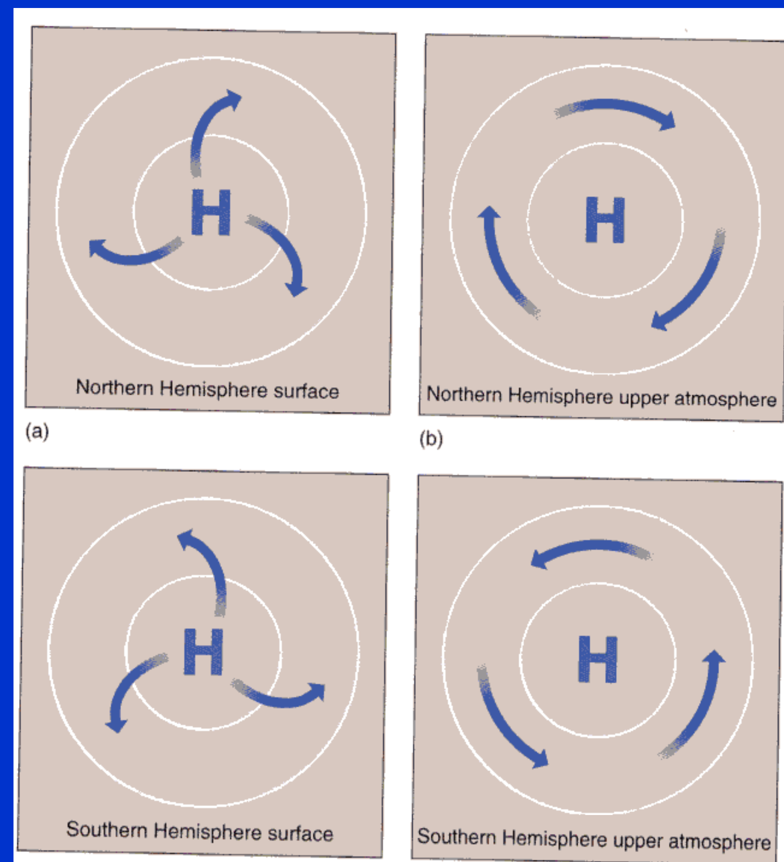
  V = wind speed

Surface Geostrophic Flow

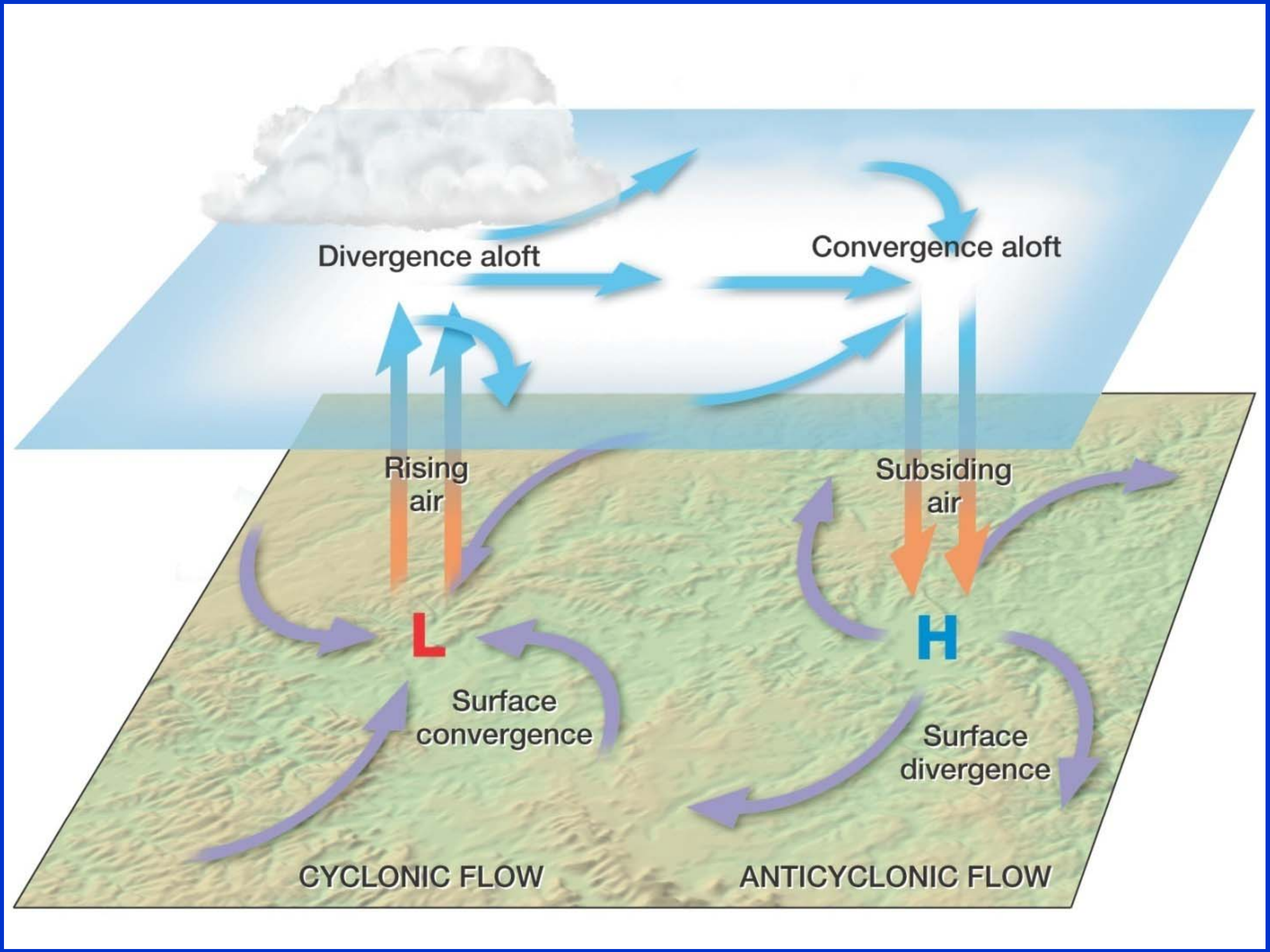
Cyclonic Flow



Anticyclonic Flow

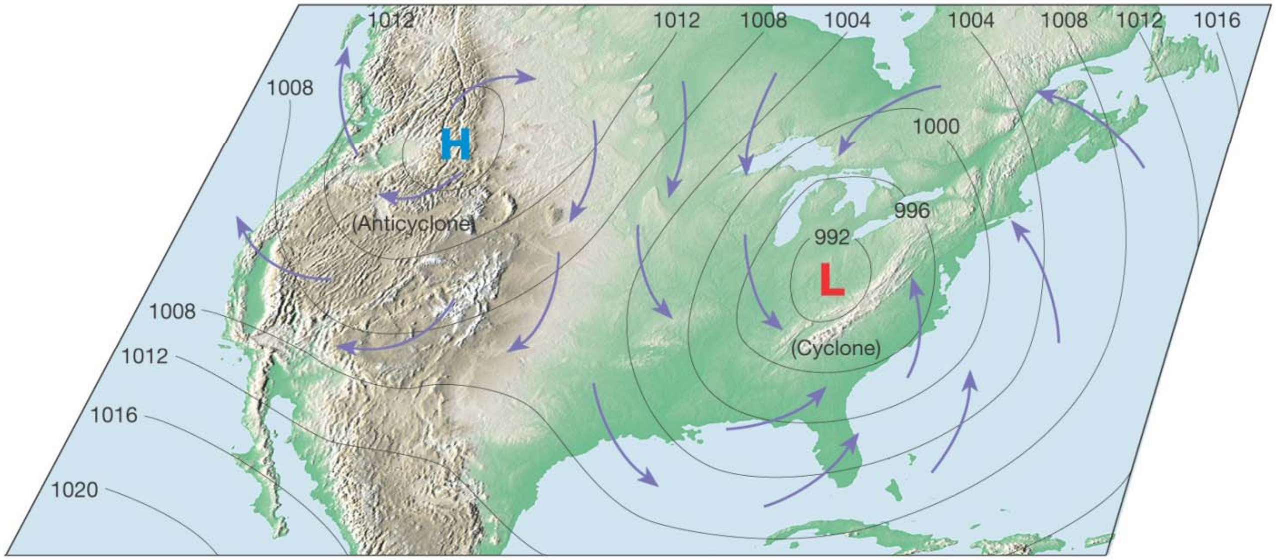


(figures from *Weather & Climate*)



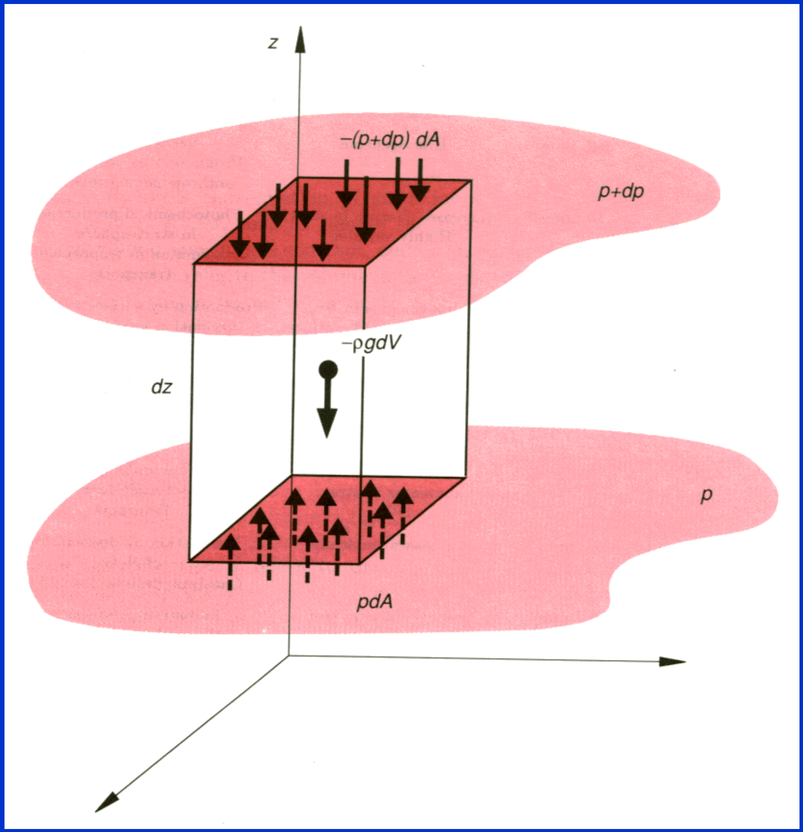
(from *The Atmosphere*)

Surface High and Low Pressure Systems



(from *The Atmosphere*)

Hydrostatic Balance in the Vertical



(from *Climate System Modeling*)

- vertical pressure force = gravitational force

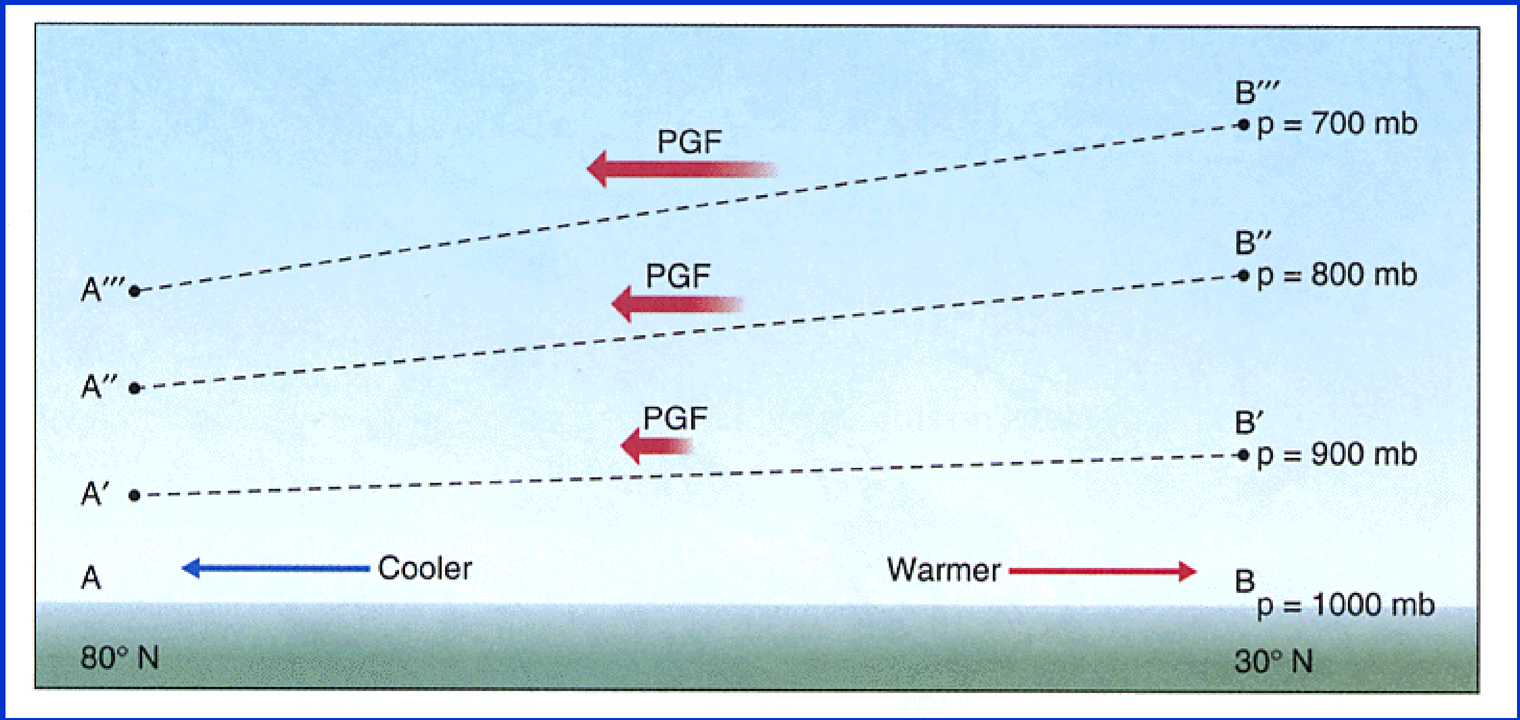
$$-(dP) \times (dA) = \rho \times (dz) \times (dA) \times g$$

$$dP = -\rho g dz$$

$$dP/dz = -\rho g$$

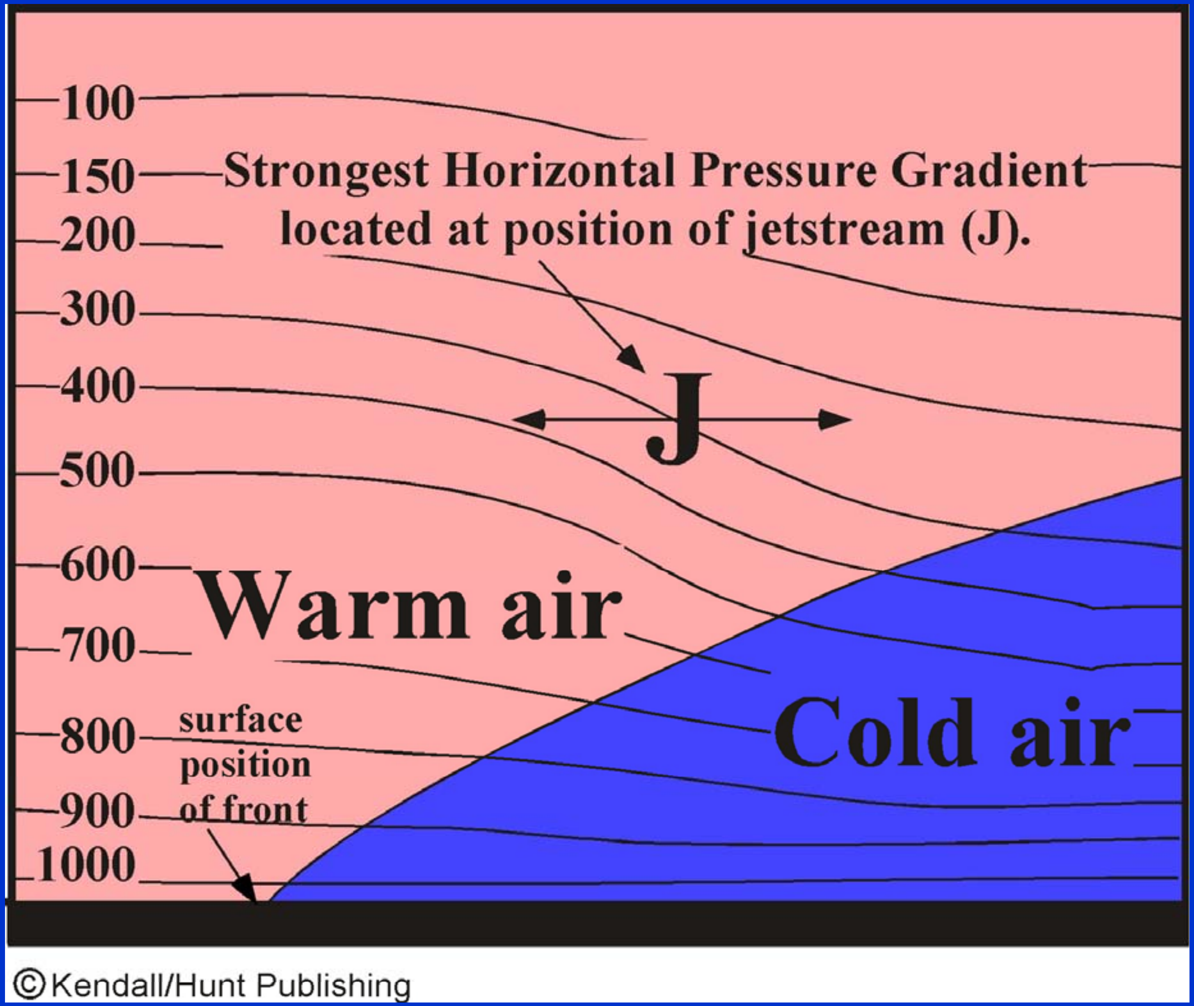
The hydrostatic balance !!

Thermal Wind Relation



(from *Weather & Climate*)

Jetstream and Front

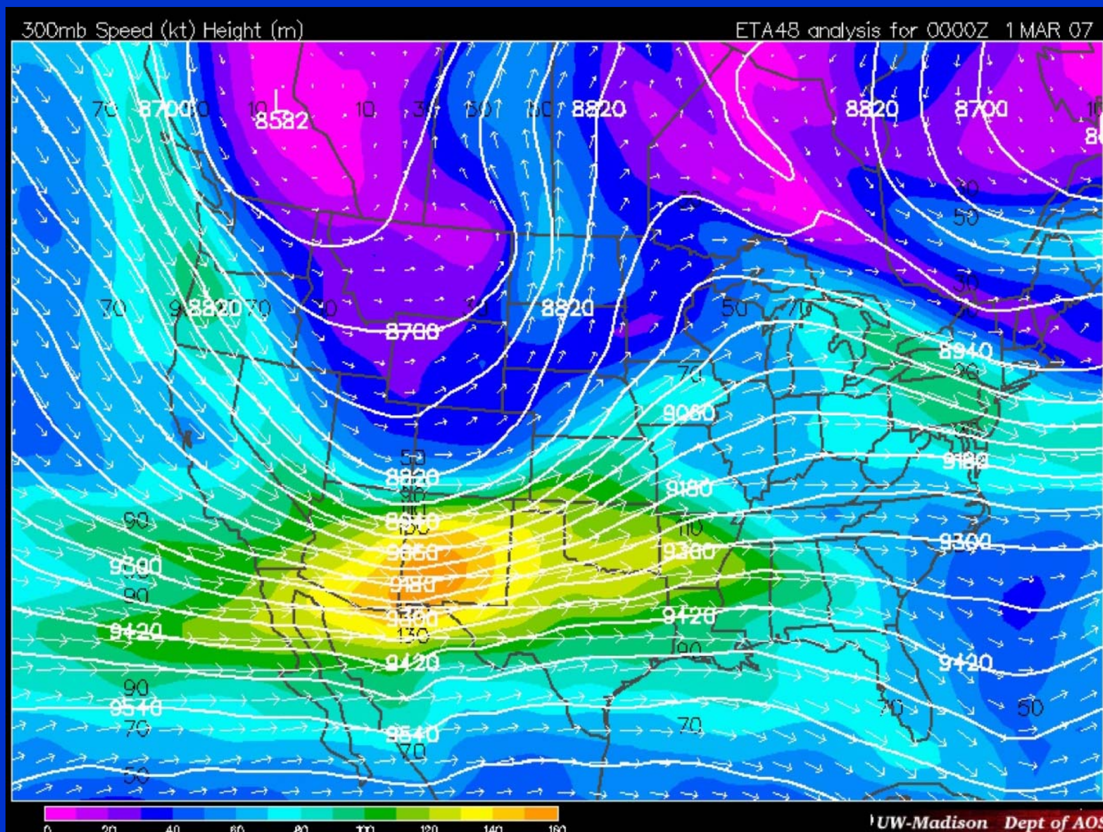


Thermal Wind Equation

$$\partial U/\partial z \;\propto\; \partial T/\partial y$$

- The vertical shear of zonal wind is related to the latitudinal gradient of temperature.
- Jet streams usually are formed above baroclinic zone (such as the polar front).

Jetstream



Courtesy of the Department of Atmospheric and Oceanic Sciences
University of Wisconsin-Madison

- The jetstream is a narrow band of strong winds that encircles the Earth in the mid-latitude.
- The band of strongest winds is typically 300 to 500 km wide and can extend from near the tropopause to about 500mb.
- The jetstream typically follows a wavelike pattern.

Three Different Jetstreams

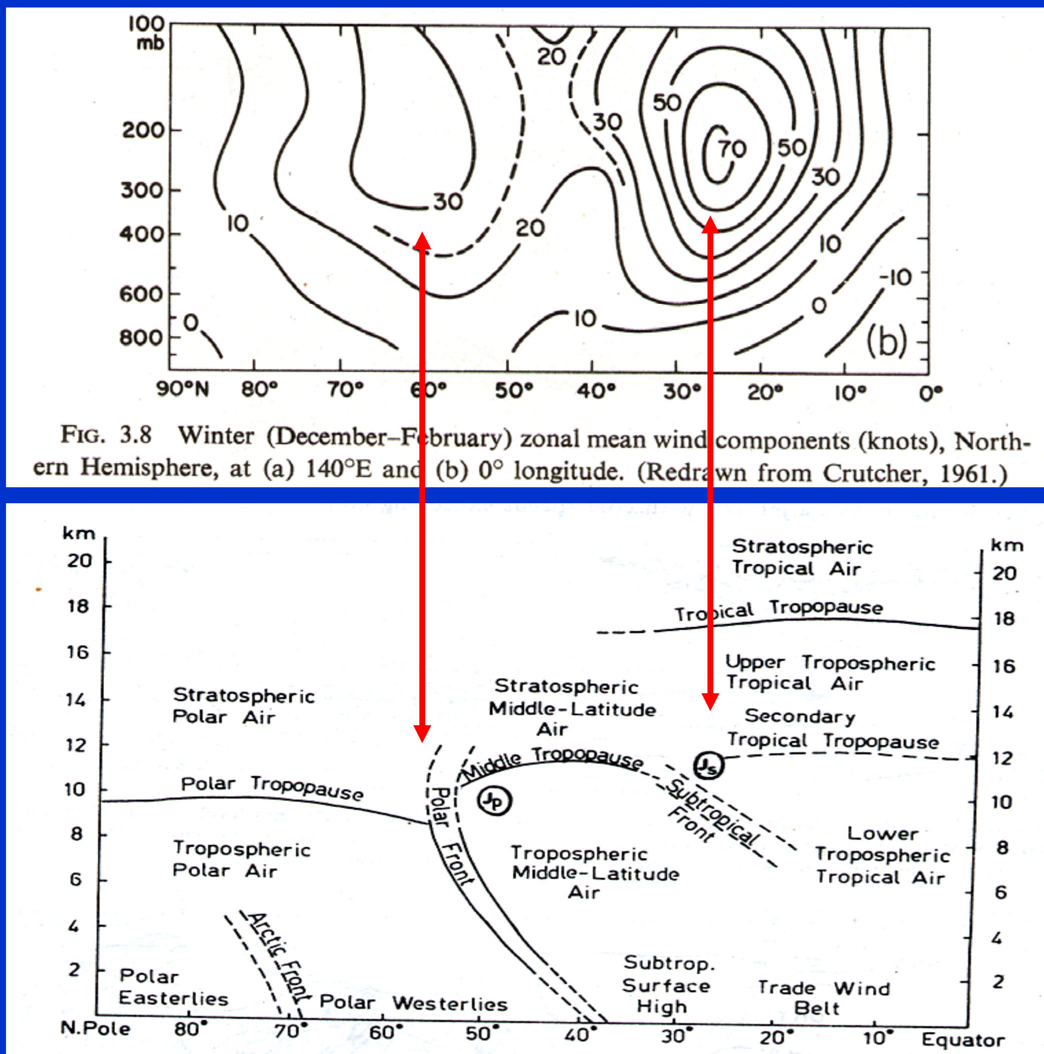


FIG. 3.8 Winter (December–February) zonal mean wind components (knots), Northern Hemisphere, at (a) 140°E and (b) 0° longitude. (Redrawn from Crutcher, 1961.)

(from *Atmospheric Circulation Systems*)

❑ Subtropical Jet

Located at the higher-latitude end of the Hadley Cell. The jet obtain its maximum wind speed (westerly) due the conservation of angular momentum.

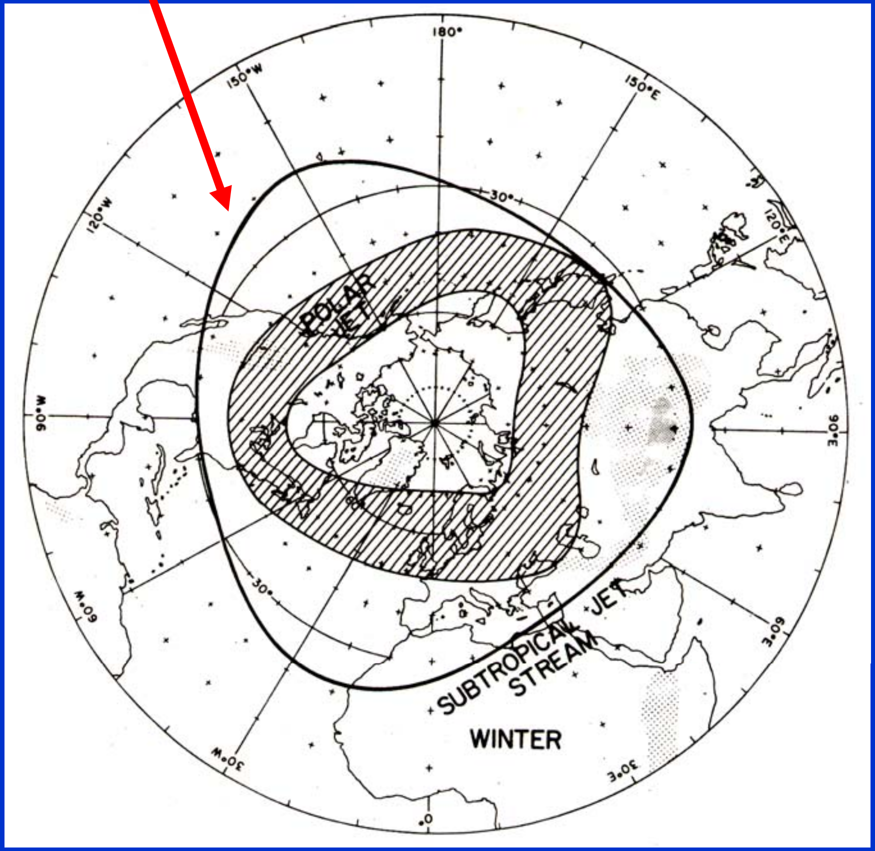
❑ Polar Jet

Located at the thermal boundary between the tropical warm air and the polar cold air. The jet obtain its maximum wind speed (westerly) due the latitudinal thermal gradient (thermal wind relation).

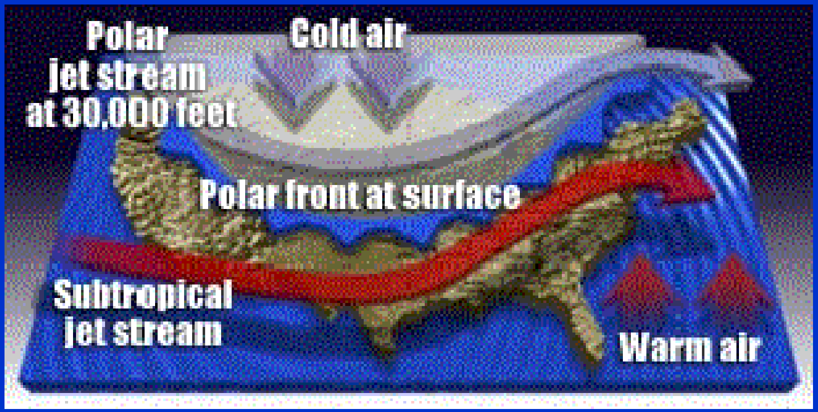
❑ Arctic Jet

Jet Streams Near the Western US

Pineapple Express



(from *Riehl (1962), Palmen and Newton (1969)*)



- ❑ Both the polar and subtropical jet streams can affect weather and climate in the western US (such as California).
- ❑ El Nino can affect western US climate by changing the locations and strengths of these two jet streams.