

Chapter 6: Stability

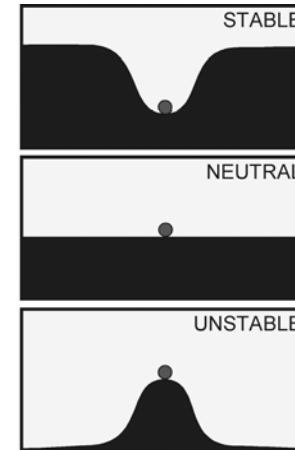


Courtesy of Glen Romine

- Concept of Stability
- Lapse Rates
- Determine Stability and Stability Indices



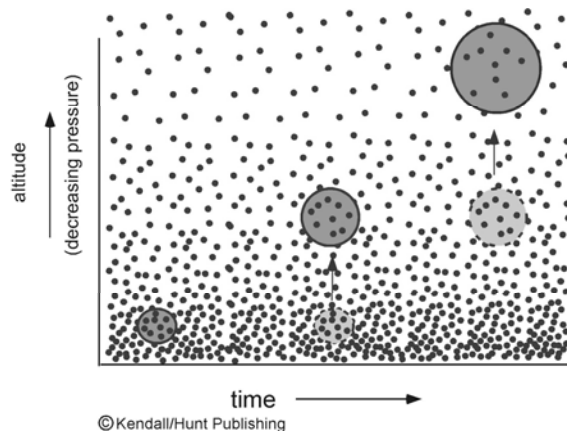
Concept of Stability



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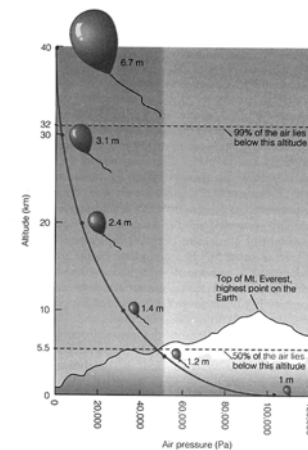
Air Parcel Expands as It Rises



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Air Parcel Expands As It Rises...



(from *The Blue Planet*)



- Air pressure decreases with elevation.
- If a helium balloon 1 m in diameter is released at sea level, it expands as it floats upward because of the pressure decrease. The balloon would be 6.7 m in diameter as a height of 40 km.

What Happens to the Temperature?

- Air molecules in the parcel (or the balloon) have to use their kinetic energy to expand the parcel/balloon.
- Therefore, the molecules lost energy and slow down their motions
 - The temperature of the air parcel (or balloon) decreases with elevation. The lost energy is used to increase the potential energy of air molecular.
- Similarly when the air parcel descends, the potential energy of air molecular is converted back to kinetic energy.
 - Air temperature rises.



Adiabatic Process

- If a material changes its state (pressure, volume, or temperature) without any heat being added to it or withdrawn from it, the change is said to be adiabatic.
- The adiabatic process often occurs when air rises or descends and is an important process in the atmosphere.

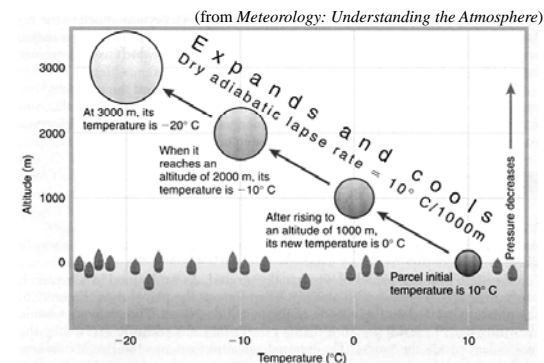


Diabatic Process

- Involve the direct addition or removal of heat energy.
- Example: Air passing over a cool surface loses energy through conduction.



Dry Adiabatic Lapse Rate



- Air parcels that do not contain cloud (are not saturated) cool at the dry adiabatic lapse rate as they rise through the atmosphere.
- Dry adiabatic lapse rate = $10^{\circ}\text{C}/1\text{km}$

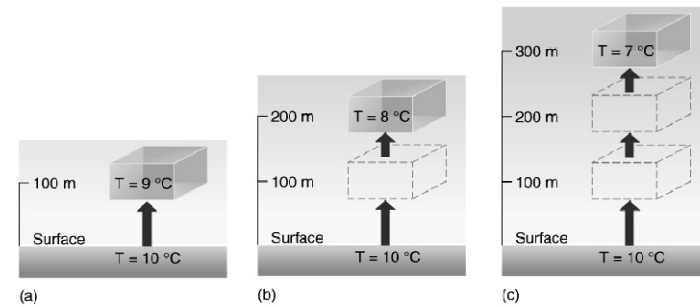


Lapse Rates

- A lapse rate is the rate at which temperature decreases (lapses) with increasing altitude.
- 3 different lapse rates we need to consider:
 - (1) dry adiabatic lapse rate
 - (2) moist adiabatic lapse rate
 - (3) environmental lapse rate

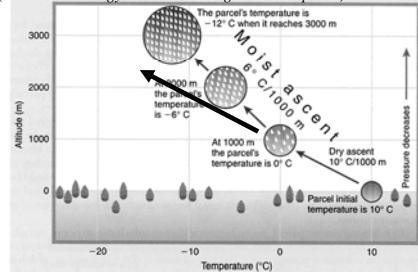


Dry Adiabatic Lapse Rate



Moist Adiabatic Lapse Rate

(from *Meteorology: Understanding the Atmosphere*)

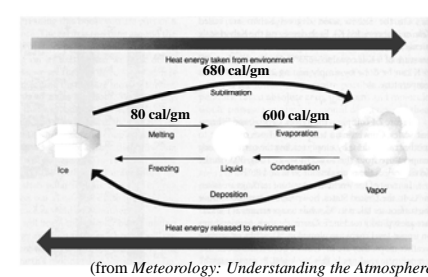


- Air parcels that get saturated as they rise will cool at a rate smaller than the dry adiabatic lapse rate due to the heating produced by the condensation of water vapor.
- This moist adiabatic lapse rate is not a constant but determined by considering the combined effects of expansion cooling and latent heating.

- In the lower troposphere, the rate is $10^\circ\text{C}/\text{km} - 4^\circ\text{C}/\text{km} = 6^\circ\text{C}/\text{km}$.
- In the middle troposphere, the rate is $10^\circ\text{C}/\text{km} - 2^\circ\text{C}/\text{km} = 8^\circ\text{C}/\text{km}$.
- Near tropopause, the rate is $10^\circ\text{C}/\text{km} - 0^\circ\text{C}/\text{km} = 10^\circ\text{C}/\text{km}$.



Phase Changes of Water



- Latent heat is the heat released or absorbed per unit mass when water changes phase.
- Latent heating is an efficient way of transferring energy globally and is an important energy source for Earth's weather and climate.

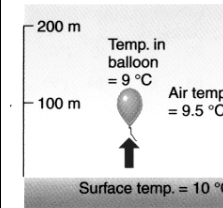


Environmental Lapse Rate

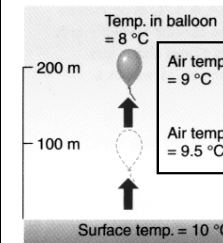
- The environmental lapse rate is referred to as the rate at which the air temperature surrounding us (or the air parcels) would be changed if we were to climb upward into the atmosphere.
- This rate varies from time to time and from place to place.
- A rawinsonde's thermometer measures the environmental lapse rate.



Environmental Lapse Rate



(a)



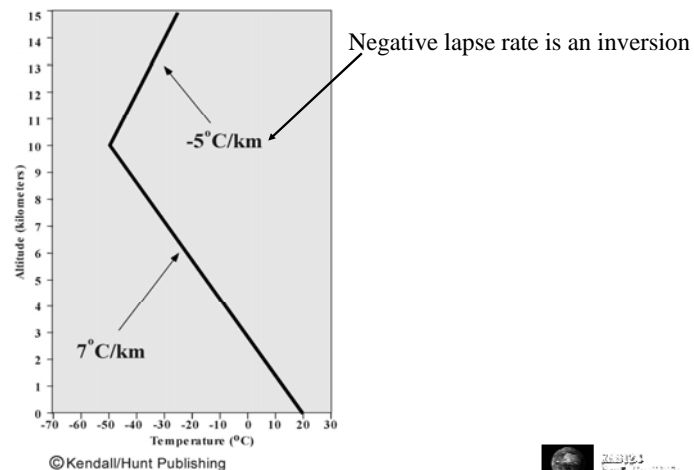
environmental lapse rate = $0.5^{\circ}\text{C}/100\text{m}$

(from *Understanding Weather & Climate*)



- The environmental (or ambient) lapse rate is referred to the vertical change in temperature through still air.
- The environmental lapse rate is not fixed. It changes from day to day and from place to place.

An Example of Environmental Lapse rate



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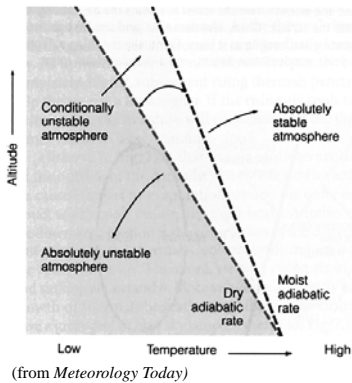


How to Determine Stability

- How do we determine where the atmosphere is unstable – under which convective clouds and storms may form?
- ➔ **Answer:** Compare the environmental lapse rate with the dry/moist lapse rate



Static Stability of the Atmosphere

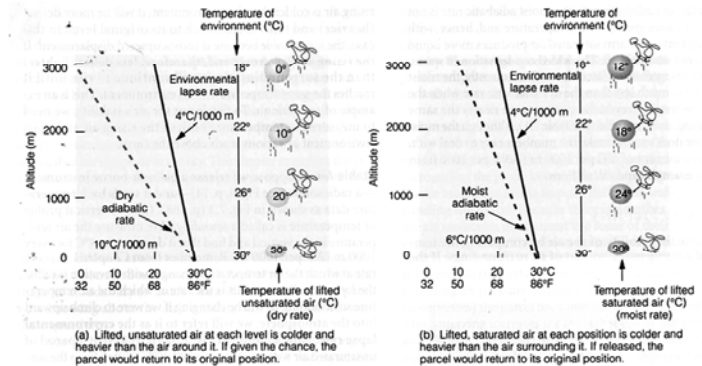


Γ_e = environmental lapse rate
 Γ_d = dry adiabatic lapse rate
 Γ_m = moist adiabatic lapse rate

- Absolutely Stable
 $\Gamma_e < \Gamma_m$
- Absolutely Unstable
 $\Gamma_e > \Gamma_d$
- Conditionally Unstable
 $\Gamma_m < \Gamma_e < \Gamma_d$



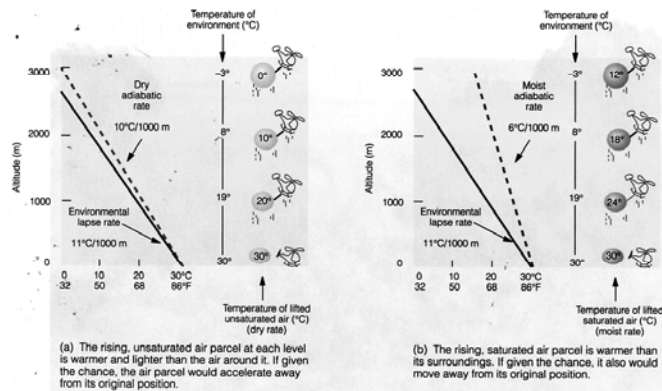
Absolutely Stable Atmosphere



(from *Meteorology Today*)



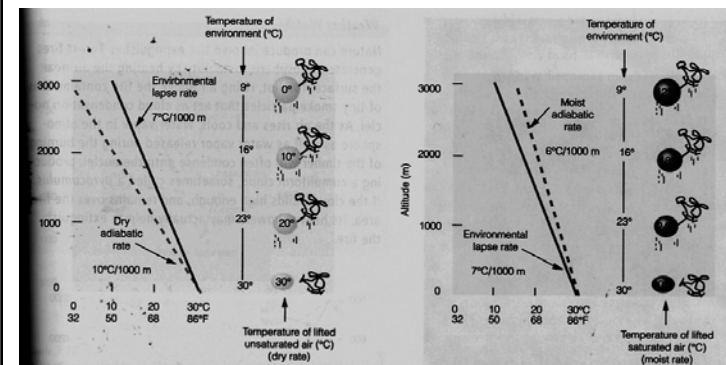
Absolutely Unstable Atmosphere



(from *Meteorology Today*)



Conditionally Unstable Atmosphere

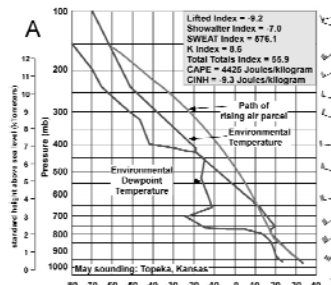


(from *Meteorology Today*)

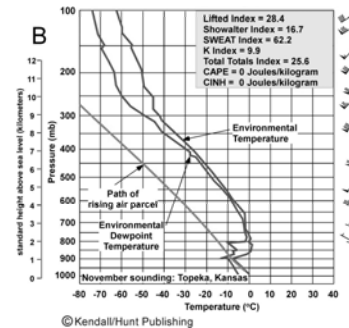


Real-Life Examples in Topeka, Kansas

Unstable Atmosphere (May)



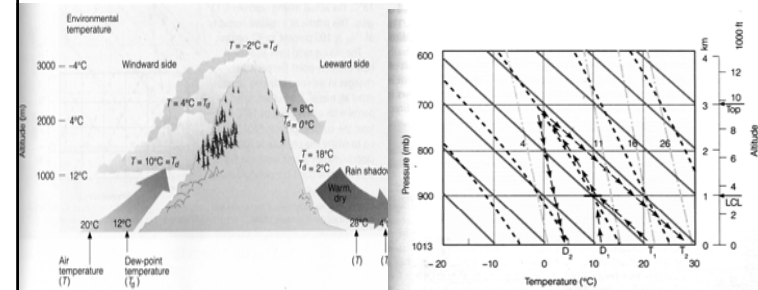
Stable Atmosphere (November)



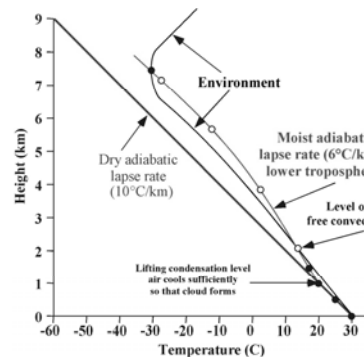
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An Example



How Thunderstorm Forms?



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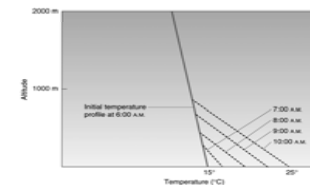
- Airs pass the level of free convection can form thunderstorms.



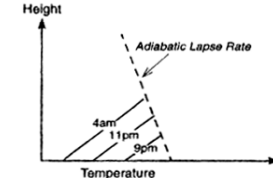
- Condensation Level**
where saturation first occurs and where cloud base is formed.
- Lifting Condensation Level**
if the air is lifted to reach the condensation.
- Level of Free Convection**
where the air first becomes buoyant (its temperature first exceeds the surrounding environment's temperature)

How to Change Environmental Lapse Rate?

Daytime



Nighttime

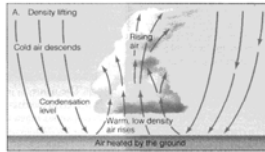


- During the day, surface insolation gains result in greater heating near the surface than aloft.
- At night, the situation reverses as terrestrial radiation loss causes near surface chilling → a *temperature inversion*.

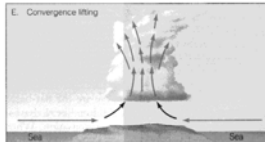


Four Ways to Lift Air Upward

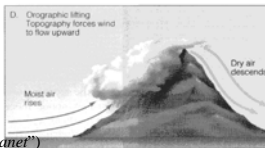
(1) Localized Convection



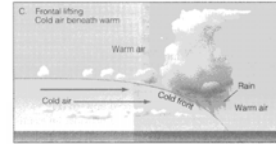
(2) Convergence Lifting



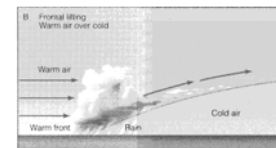
(3) Orographic Lifting



(from "The Blue Planet")



(4) Frontal Lifting



Stability Indices

- (1) Environmental Lapse rate
- (2) Lifted Index = $T(\text{environment at } 500\text{mb}) - T(\text{parcel lifted to } 500\text{mb})$
- (3) Showalter Index: similar to lifted index but was lifted to 850mb
- (4) CAPE (Convective Available Potential Energy): derived from soundings
- (5) Convective INHibition (CINH) Index
- (6) K Index
- (7) Total Totals Index
- (8) SWEAT (Severe Weather Threat) Index

TABLE 6.2

Stability Categories and Likelihood of Severe Convective Storms for Various Ranges of the Lifted Index (LI), Showalter Index (SI), Convective Available Potential Energy (CAPE), Total Totals (TT) Index and SWEAT (SW) Index

Stability	LI	SI	CAPE	TT	SW
Very stable (no significant activity)	> +3				
Stable (Showers possible; T's showers unlikely)	0 to +3	> +2	< 0		
Marginally unstable (T's showers possible)	-2 to 0	0 to 2	0 to 1000	45 to 50	
Moderately unstable (Thunderstorms possible)	-4 to -2	-3 to 0	1000 to 2500	50 to 55	250 to 300
Very unstable (Severe T's storms possible)	-6 to -4	-6 to -3	2500 to 3500	55 to 60	300 to 400
Extremely unstable (Severe T's storms probable; tornadoes possible)	< -6	< -6	> 3500	60	> 400